

## **Supplementary Material for:**

### **Late Triassic to Early Jurassic carbon isotope chemostratigraphy and organofacies evolution in a distal to proximal transect of the North German Basin**

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## **Content:**

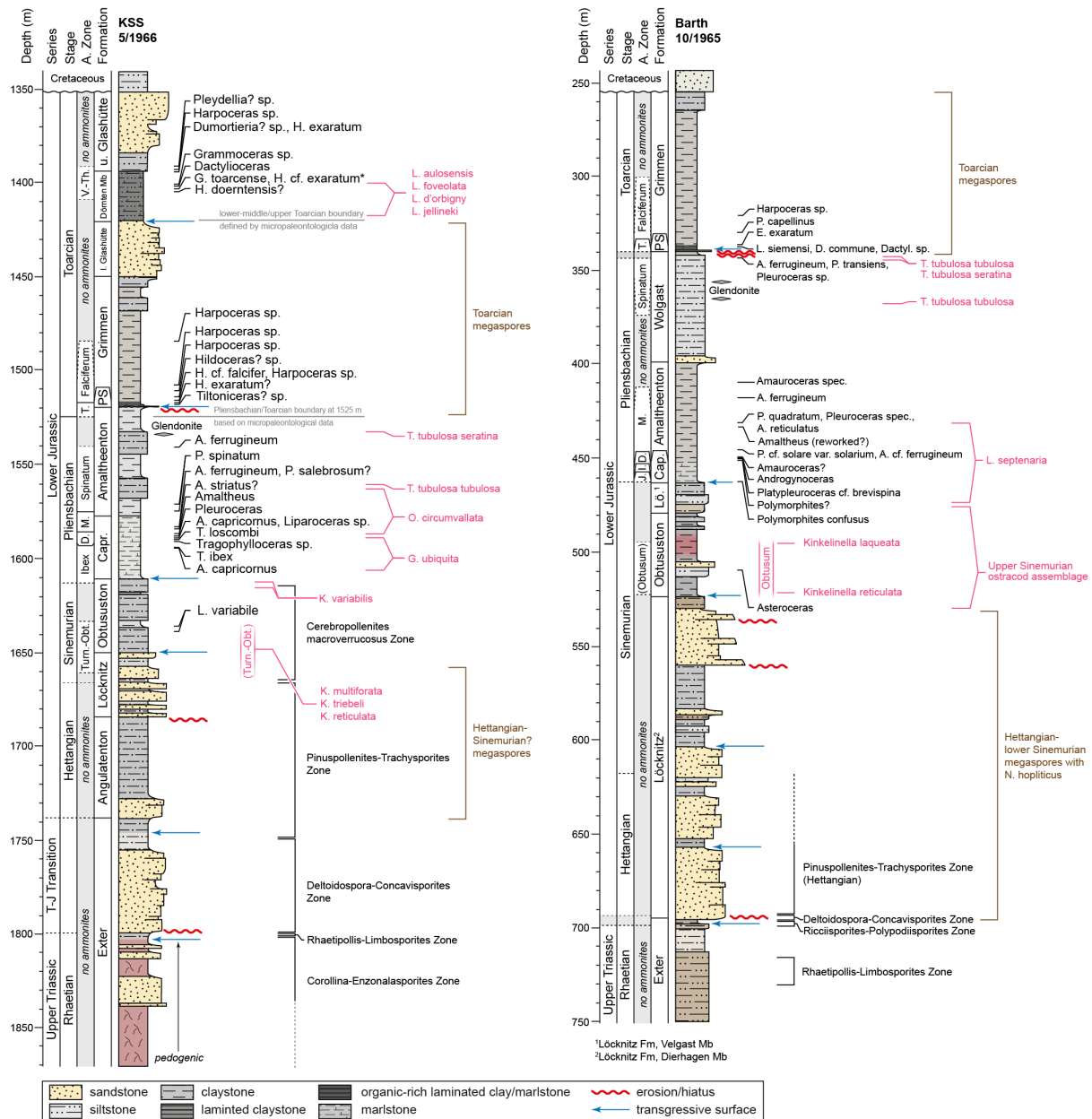
**Figure 1.** Biostratigraphic data for cores KSS 5/66 and Barth 10/65.

**Figure 2.** Chemostratigraphic correlation of Core Schandelah and the Mochras Core.

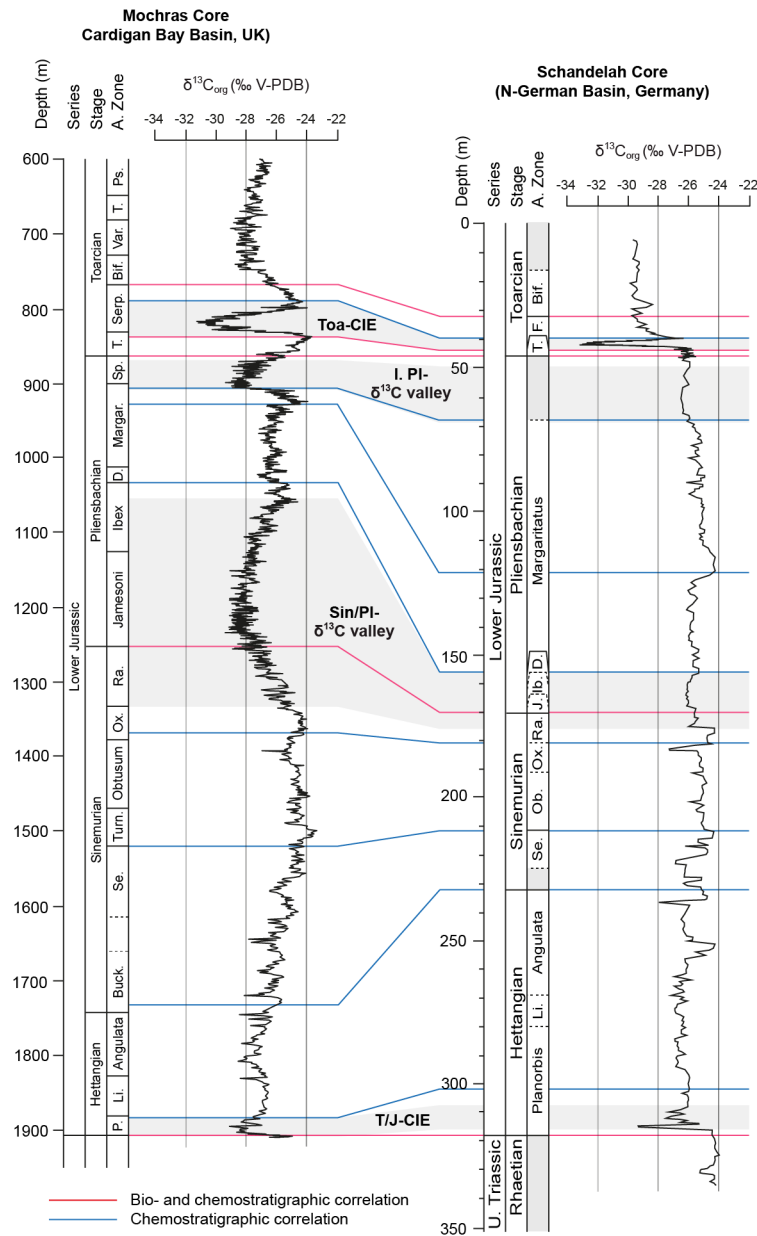
**Figures 3-5.** Cyclostratigraphic analysis for cores KSS 5/66 and Barth 10/65 (Toarcian only).

**Figure S6.** Comparison of sediment thickness along the basin-margin transect.

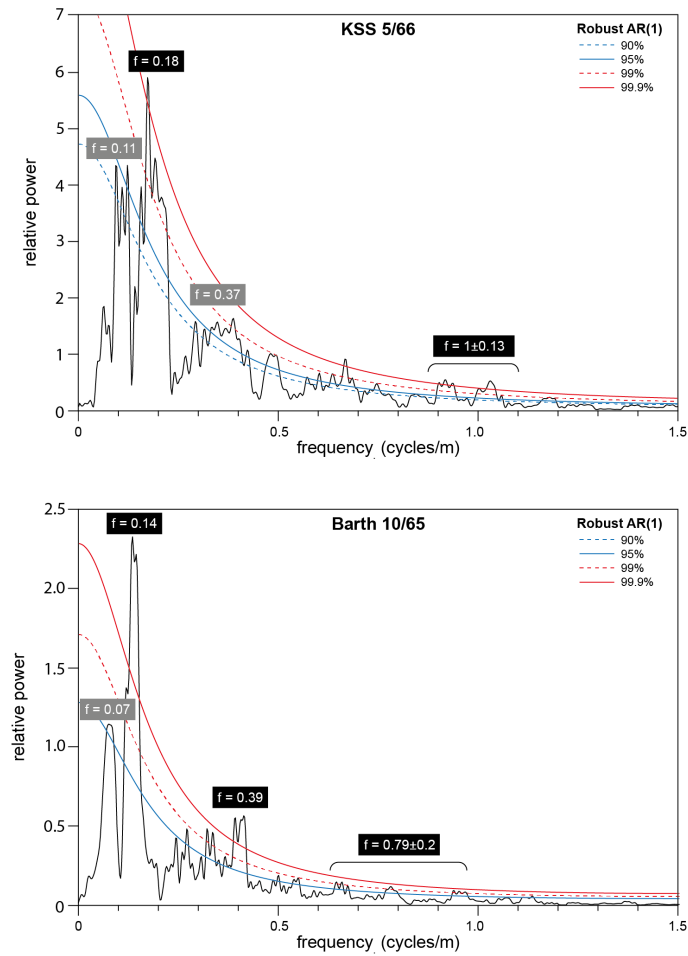
## **Methods – cyclostratigraphic analysis**



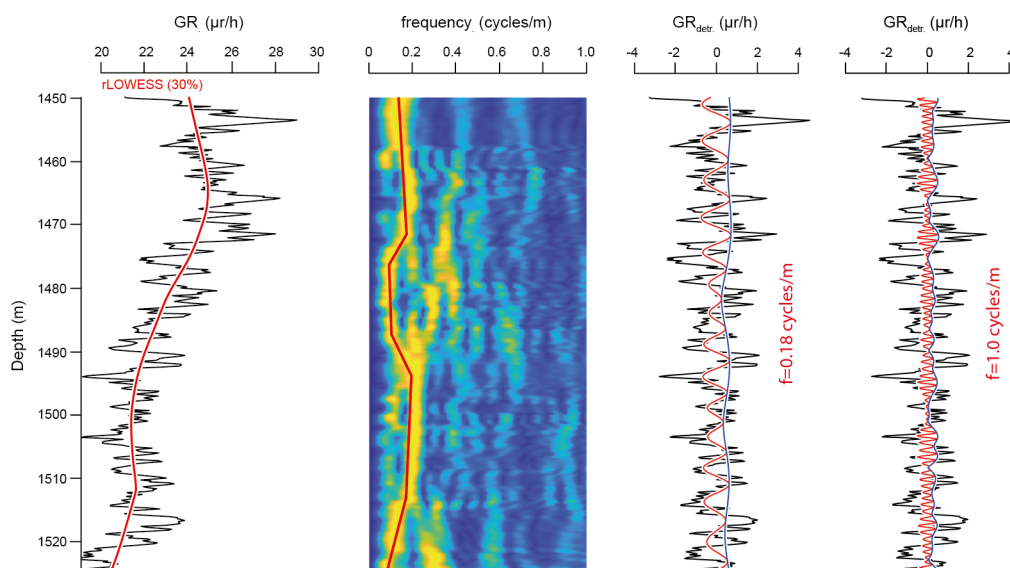
**Fig. S1:** Biostratigraphic framework for cores KSS 5/1966 and Barth 10/1965 (Zimmermann et al. 2015, 2018; Barth et al. 2018a, 2018b and references therein) (abbreviations of ammonite zones: Turn.: Turneri, Obt.: Obtusum, D.: Davoei, I: Ibex; J: Jamesoni; M.: Margaritatus, T.: Tenuicostatum, V.: Variabilis, Th.: Thouarensis; abbreviations for formations: Cap.: Capricornumergel, PS: Posidonienschiefer; l./u. Glashütte: lower/upper Glashütte).



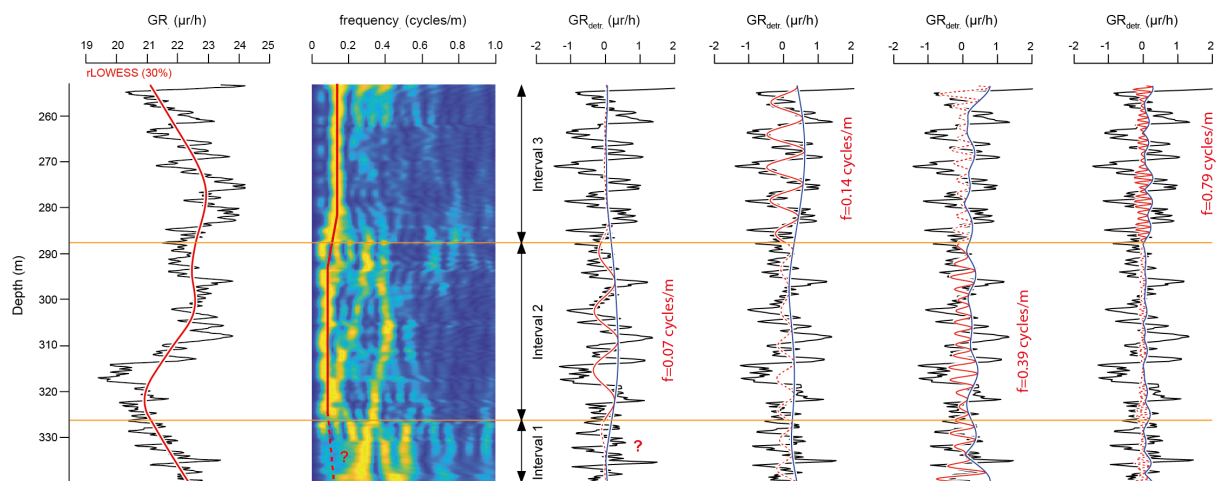
**Fig. S2:** Biostratigraphic and chemostratigraphic correlation of the Mochras Core (Xu et al. 2018; Storm et al. 2020) with Core Schandelah (van de Schootbrugge et al. 2018) (Buck.: Bif.: Bifrons; Bucklandi; D.: Davoei; Li.: Liasicus; Margar.: Margaritatus; Ox.: Oxynotum; P.: Planorbis; Ps.: Pseudoradosa; Ra.: Raricostatum; Sp.: Serp.: Serpentinum; Sp.: Spinatum; T.: Tenuicostatum; Turn.: Turner; Var.: Variabilis). Both cores reveal comparable trends in sedimentary  $\delta^{13}C_{org}$  values. Differences are seen in the magnitude of isotopic changes, as well as in the magnitude of  $\delta^{13}C_{org}$  absolute values. In particular, the  $\delta^{13}C$  valley across the Sinemurian/Pliensbachian boundary (Sin/PI  $\delta^{13}C$  valley) and in the late Pliensbachian (I. PI  $\delta^{13}C$  valley) show lower values in the Mochras Core (T/J-CIE: Triassic/Jurassic carbon isotope excursion; Toa-CIE: Toarcian carbon isotope excursion). This indicates that trend in  $\delta^{13}C_{org}$  values reflect trans-regional (global) carbon cycle dynamics. Absolute values in the  $\delta^{13}C_{org}$  values, however, reflect local processes.



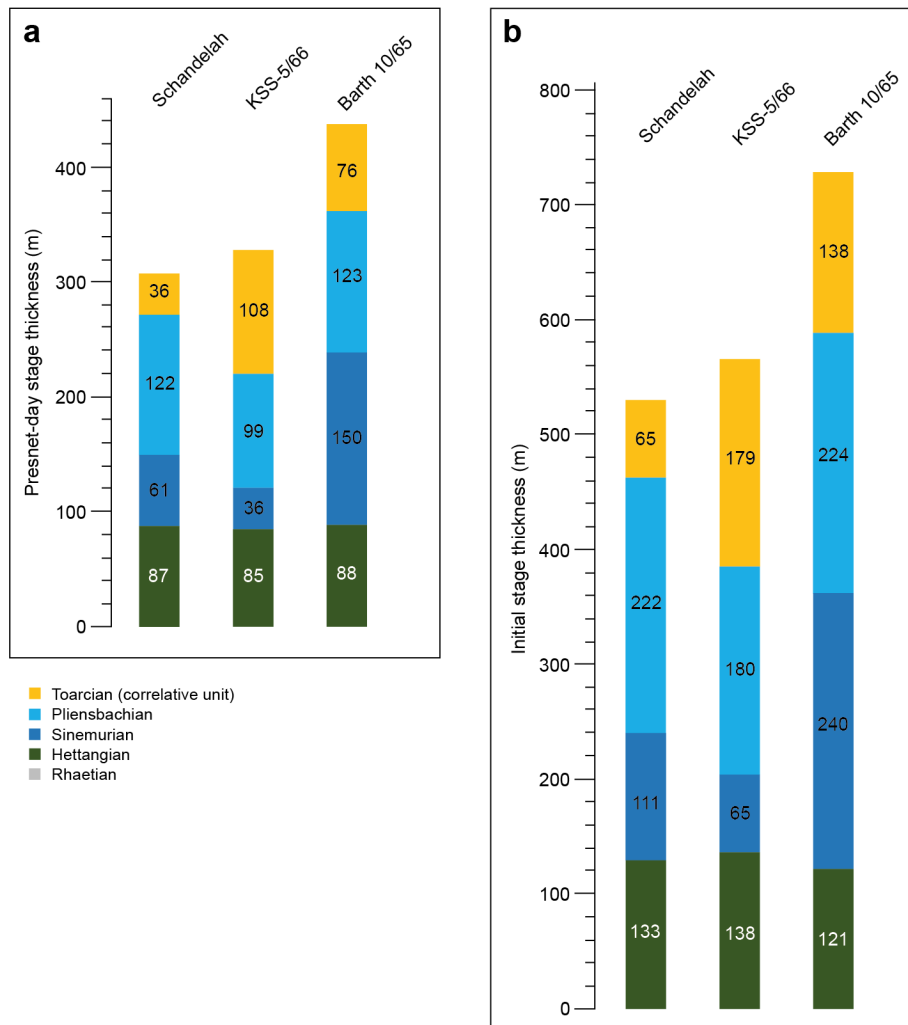
**Fig. S3:** MTM power spectra for gamma ray data of Toarcian strata in Core KSS 5/66 (top) and Barth 10/65 (bottom)



**Fig. S4:** Gamma ray data, evolutionary power spectra and filter outputs for Core KSS 5/66.



**Fig. S5:** Gamma ray data, evolutionary power spectra and filter outputs for Core Barth 10/65.



**Fig. S6:** Comparison of the present-day sediment thickness (**a**) and the initial stage thickness before compaction (**b**) along the distal-proximal transect. Sediment compaction and thickness reduction is the product of lithology and burial depth. Based on the low to moderate  $T_{\max}$  values we estimated the burial depth to 1000-1500 m. We used compactions factors for sandstones and mudstones of 0.55 and 0.8, respectively (Lee, 2018). However, substantially higher compaction rates of 0.1 (or lower) have been assumed for the bituminous Posidonienschiefer of southern Germany (Bandel and Knitter 1986).

## Methods – cyclostratigraphic analysis

Cyclostratigraphic investigations for the Toarcian Grimmer Formation were performed on gamma ray data. Prior to cyclostratigraphic analysis, data were detrended using the weighted-average rLOESS method. Cyclostratigraphic analyses of the carbonate and TOC data were performed using the ACYCLE 2.3 tool for Matlab (Li et al. 2019). For spectral analyses we used the  $2\pi$  multitaper method (MTM) (Thomsen 1982), together with robust red-noise models (Mann and Lees 1996). Statistically significant spectral peaks were filtered from the data series using Taner-Hilbert filtering. Evolutionary spectral analysis was performed using Fast Fourier transform (LAH) as implemented in the ACYCLE 2.3 tool for Matlab (Li et al. 2019). We used a sliding window size of 200 and step size of 2. Evolutionary spectral analysis allows assessing the evolution of spectral components along the investigated section.

We here do not aim to establish an astronomical timescale for the Posidonienschiefer and Grimmer Formations, as sedimentation will have been affected by hiatuses, condensation and varying sediment accumulation rates. We here use cyclostratigraphic analysis as additional correlation tool for the biostratigraphically-poor constrained Grimmer Formation.

## References:

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