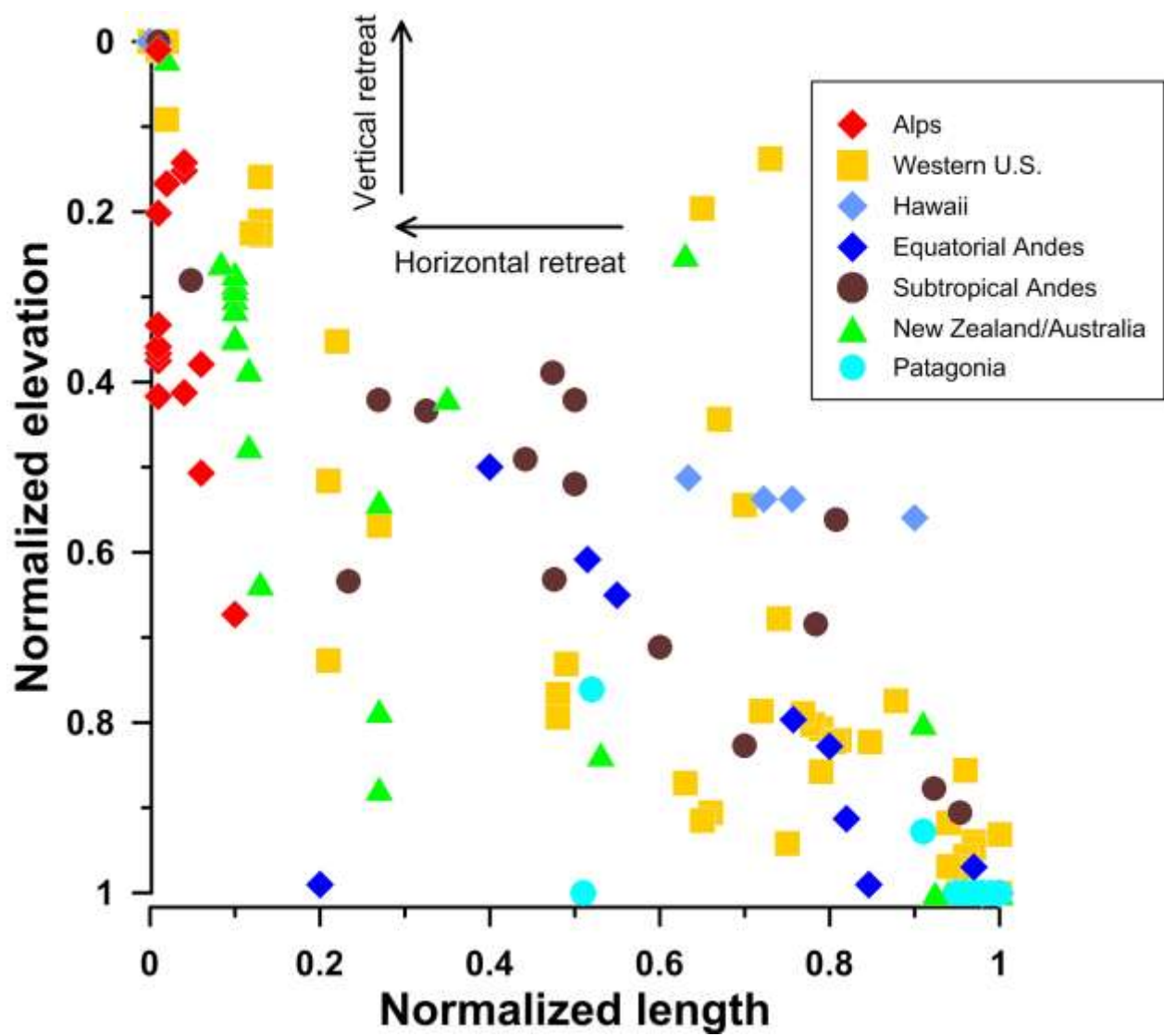
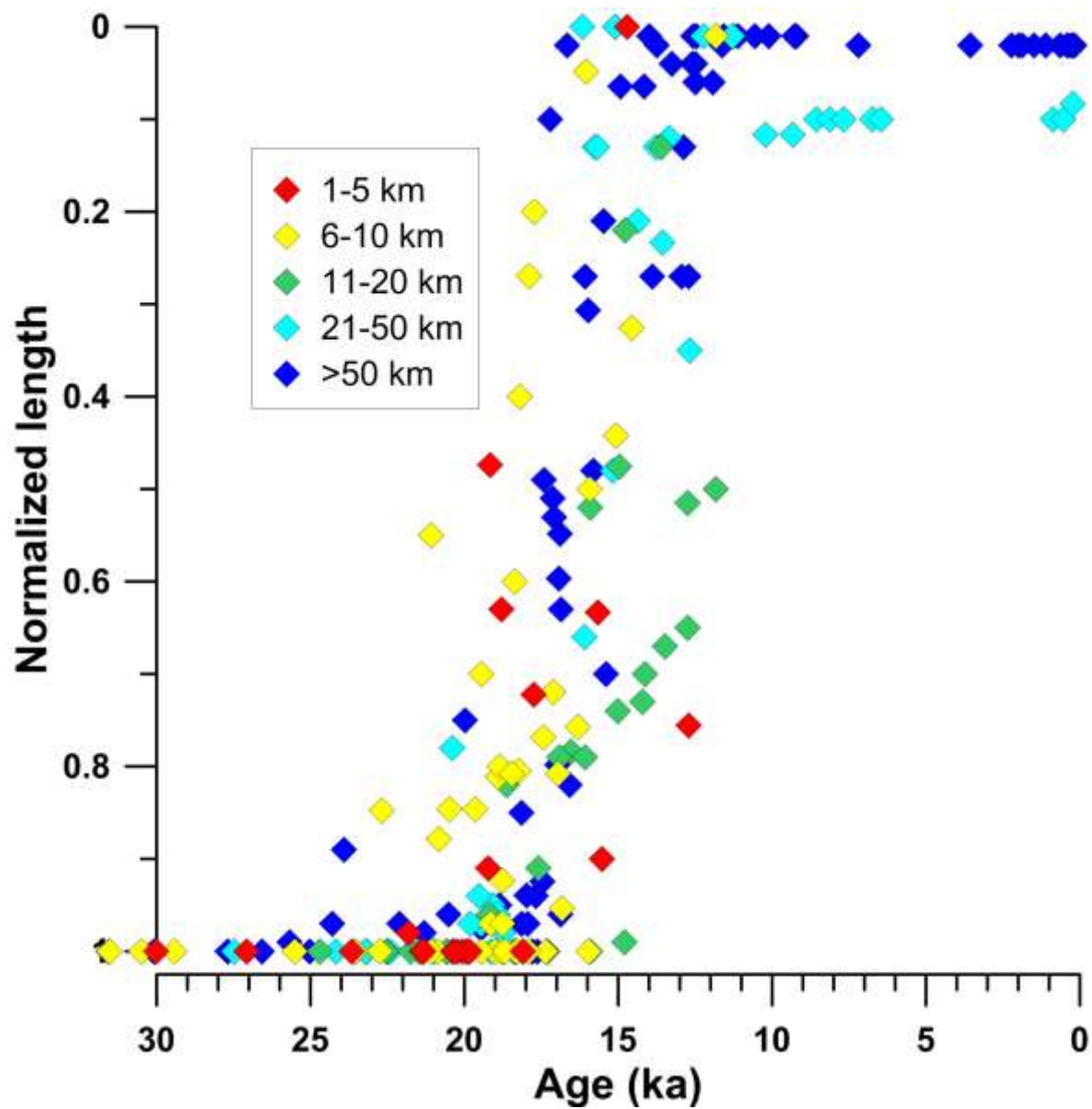
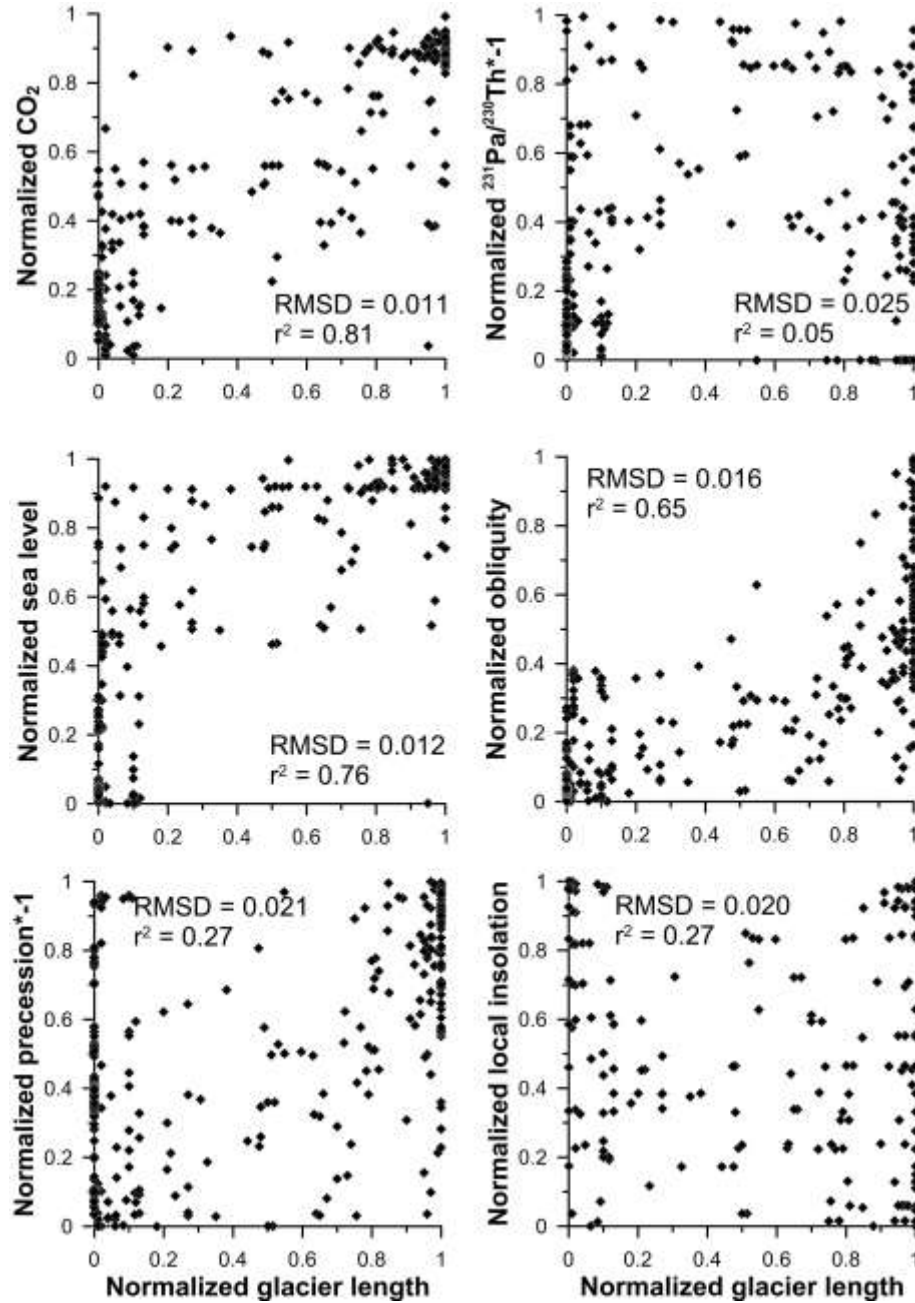




Supplementary Figure 1. Normalized glacier terminus elevation changes versus normalized glacier length changes.

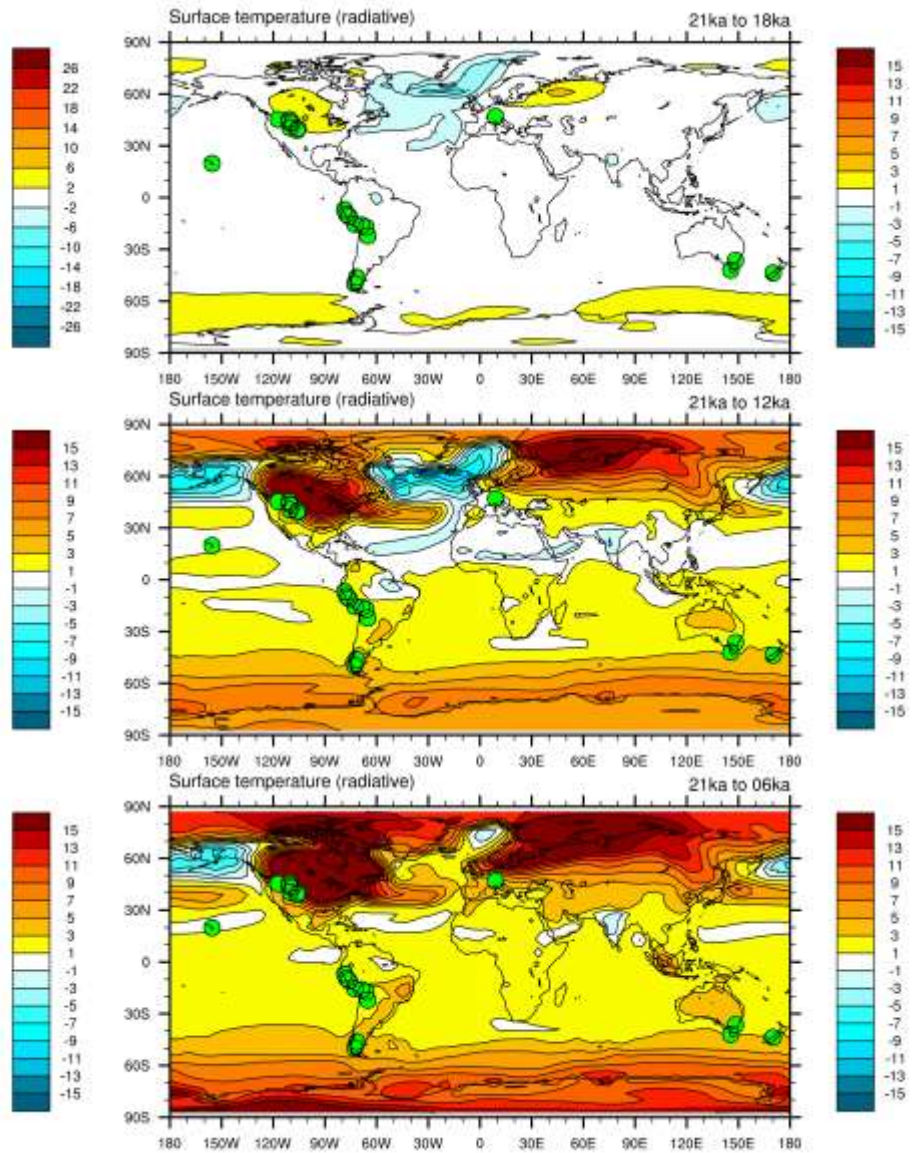
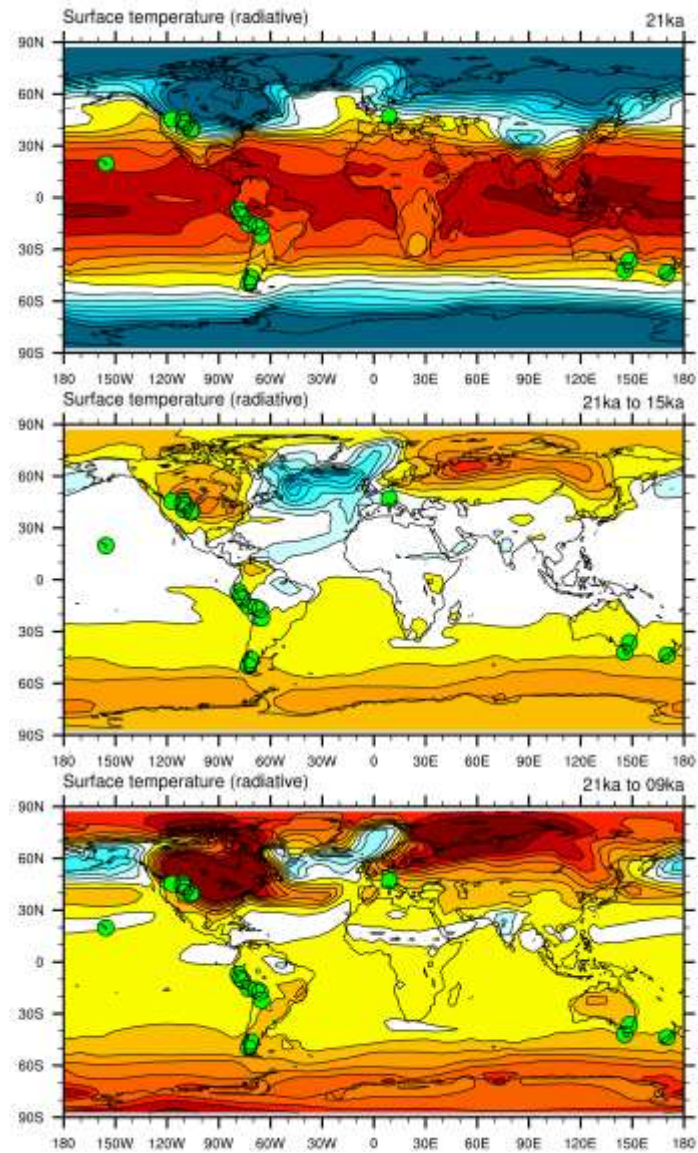


Supplementary Figure 2. Normalized glacier lengths through time, as shown in Fig. 2 in the main text, but divided into several size classes based upon their LGM extent.

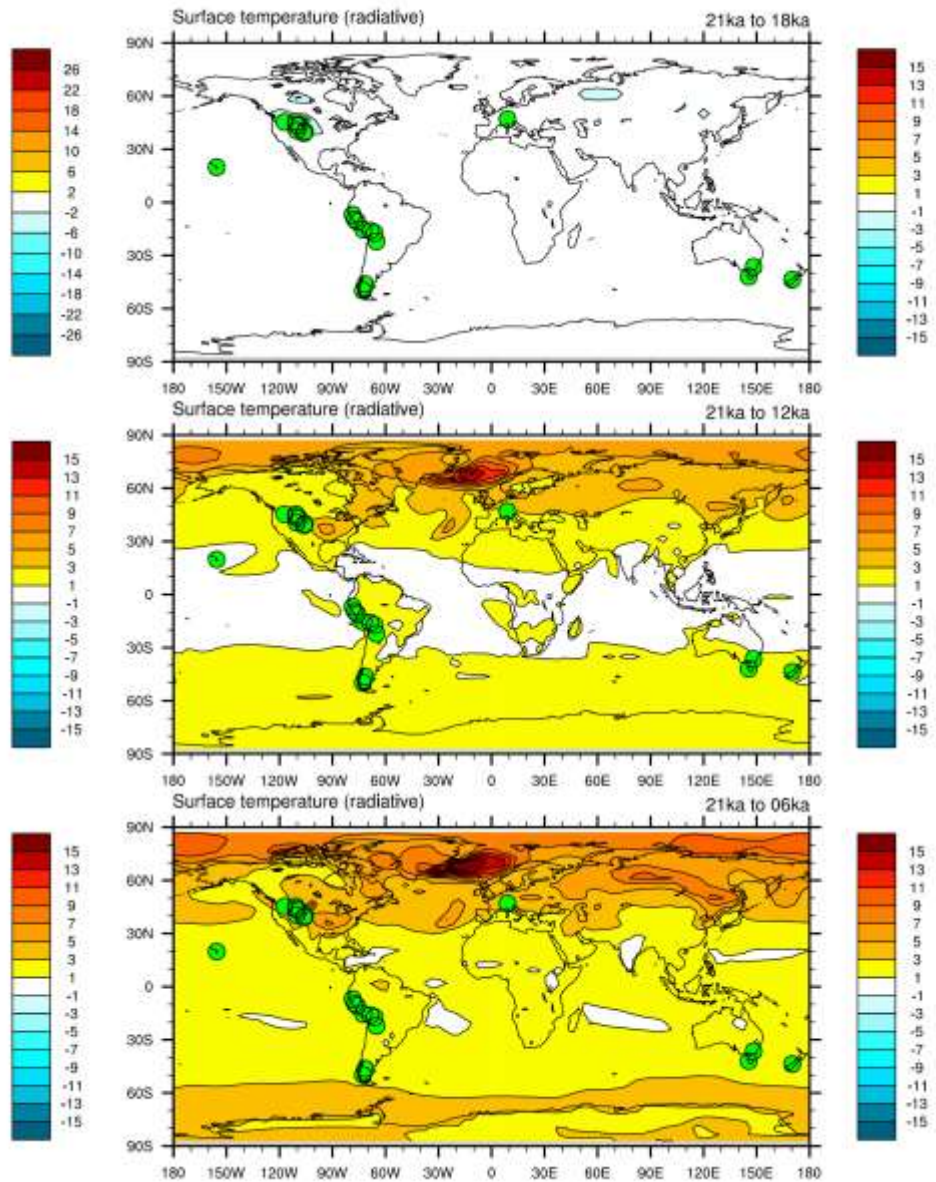
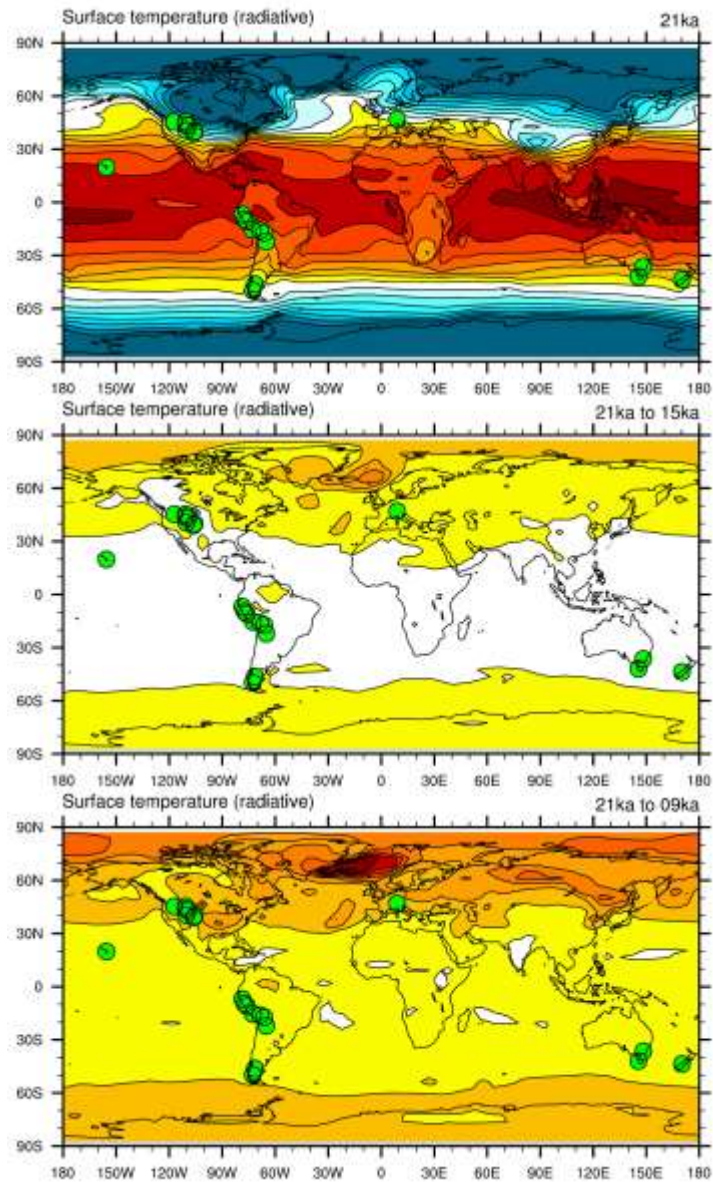


Supplementary Figure 3. Scatterplots of normalized glacier lengths (x-axes) versus normalized forcings (y-axes) showing root-mean-squared deviation (RMSD) and coefficient of determination (r^2). Forcings considered include CO_2 (ref. ¹), AMOC variations as represented by a Pa/Th record², ice sheets as represented by eustatic sea level³, obliquity, precession, and local insolation (JJA for northern extratropical sites, DJF for southern extratropical sites, mean annual for tropical sites).

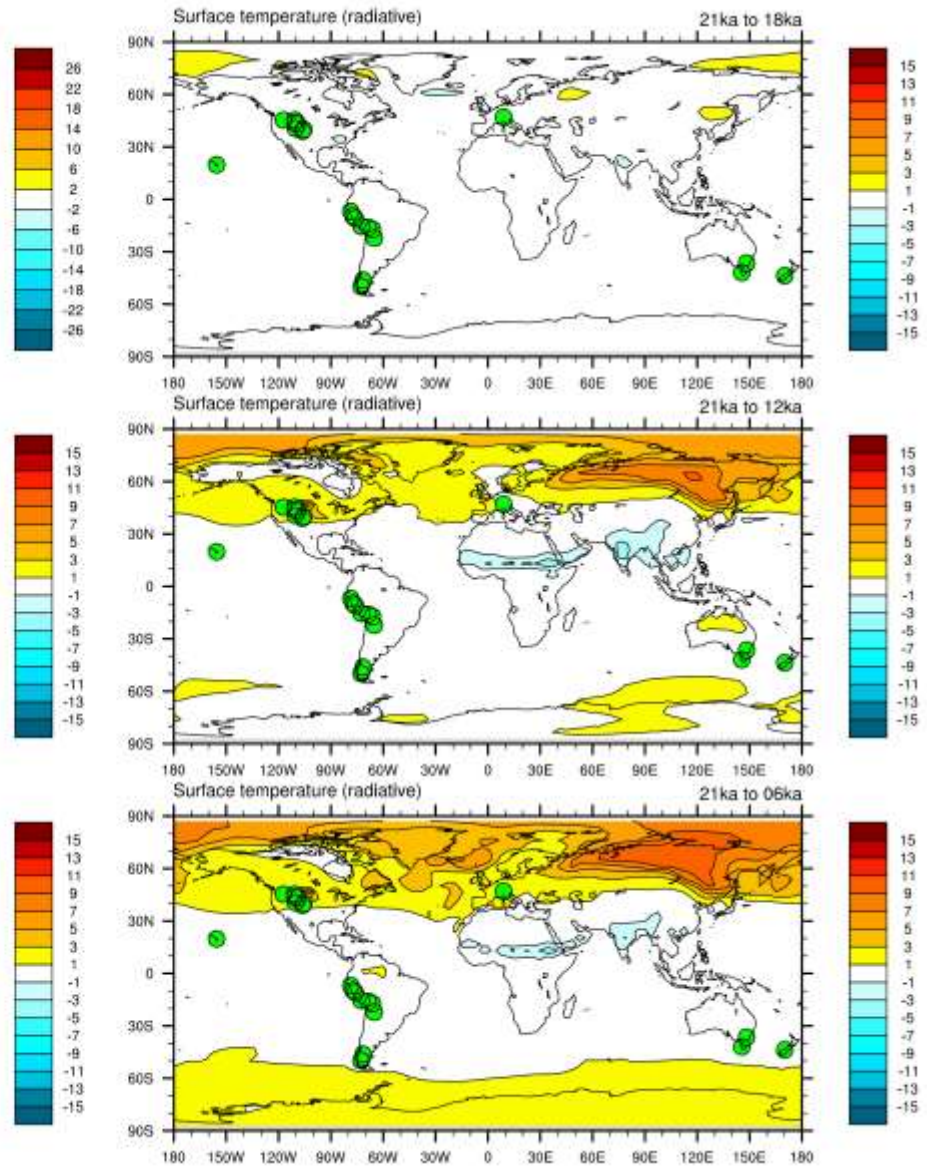
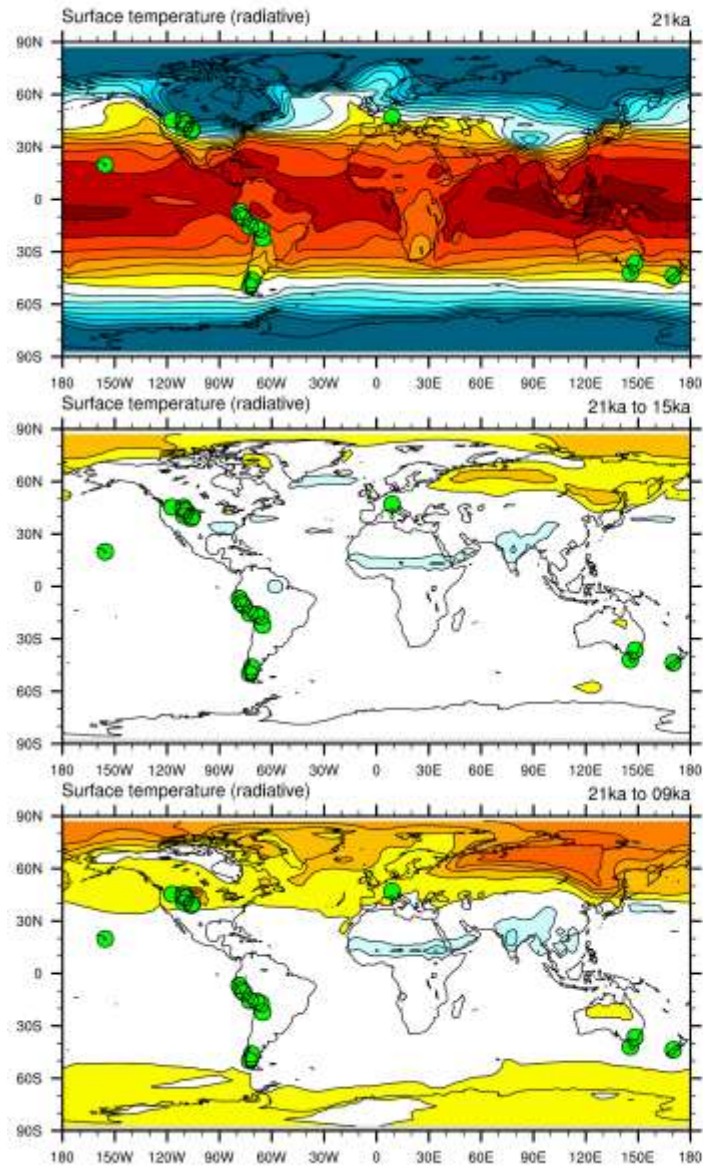
ALL simulation



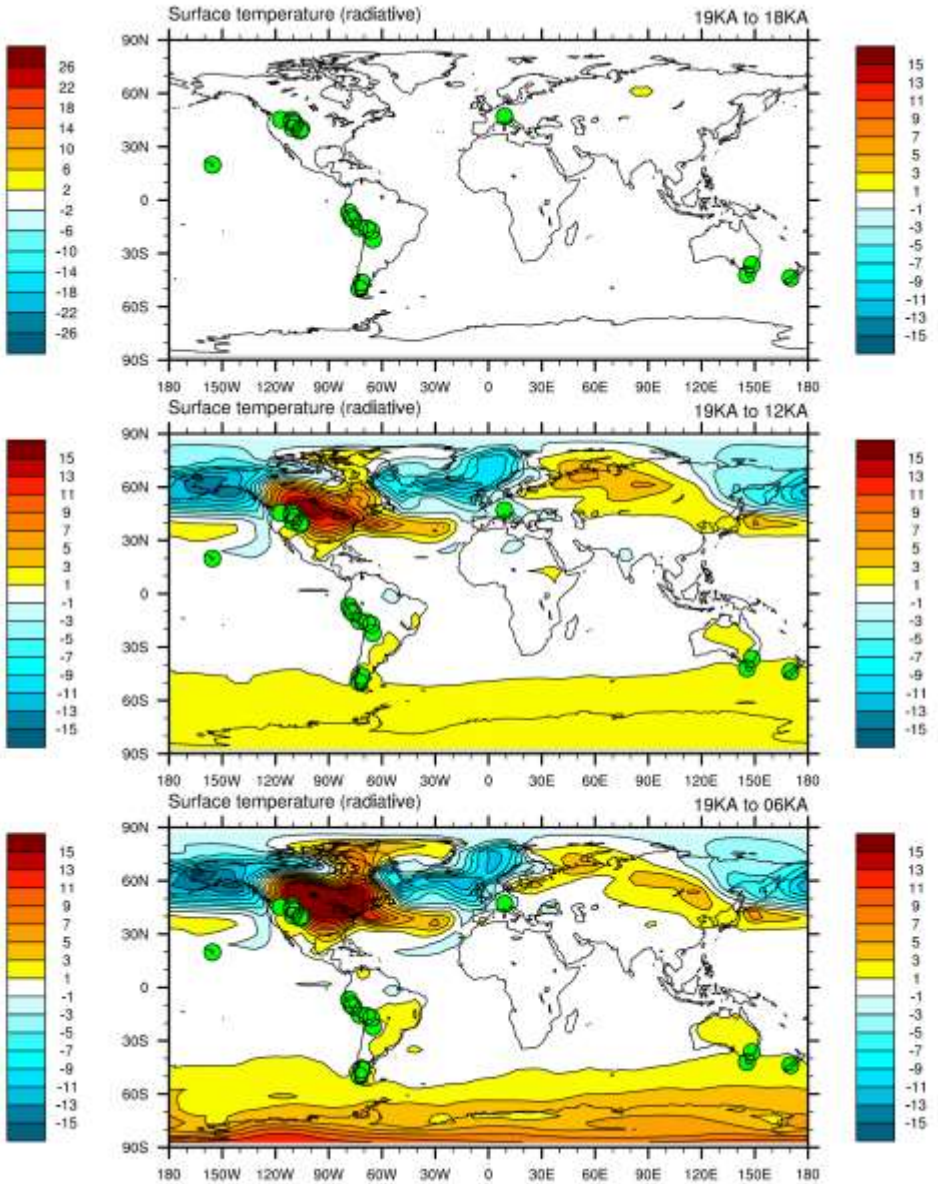
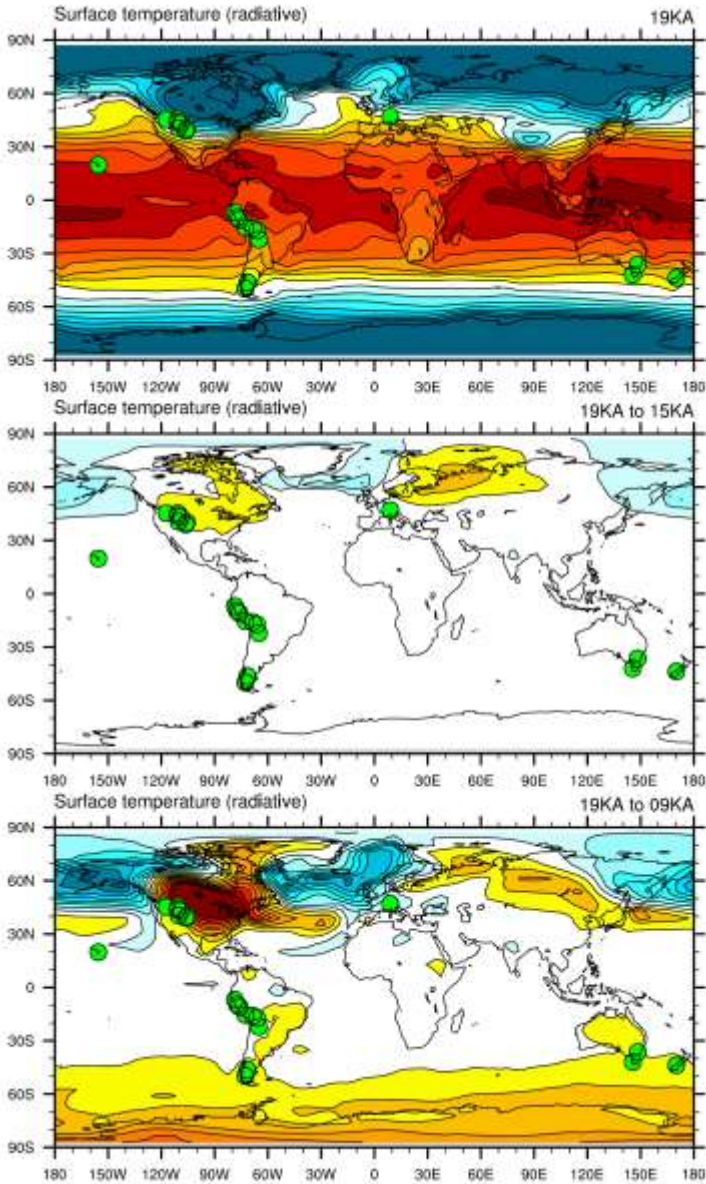
GHG simulation



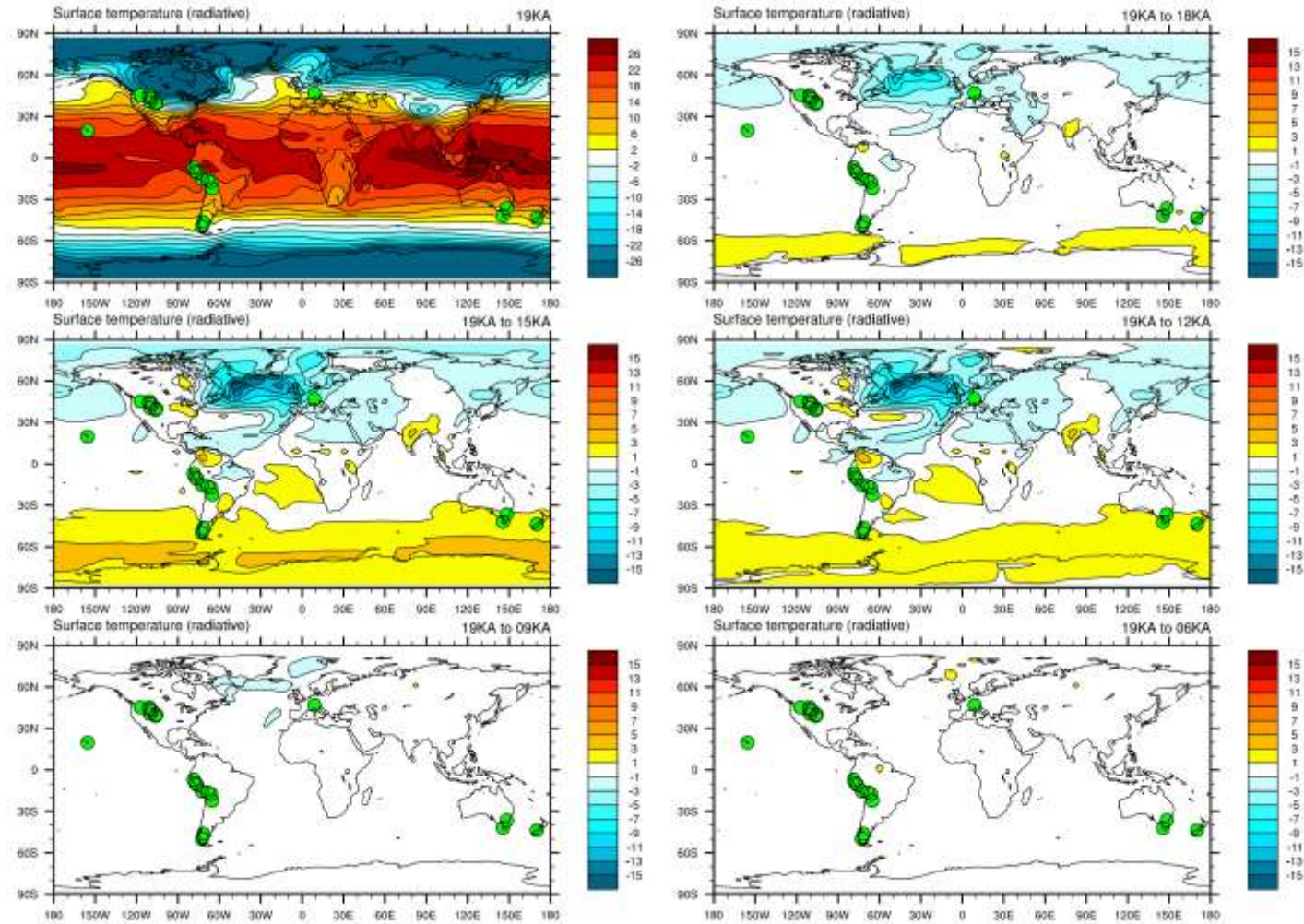
ORB simulation



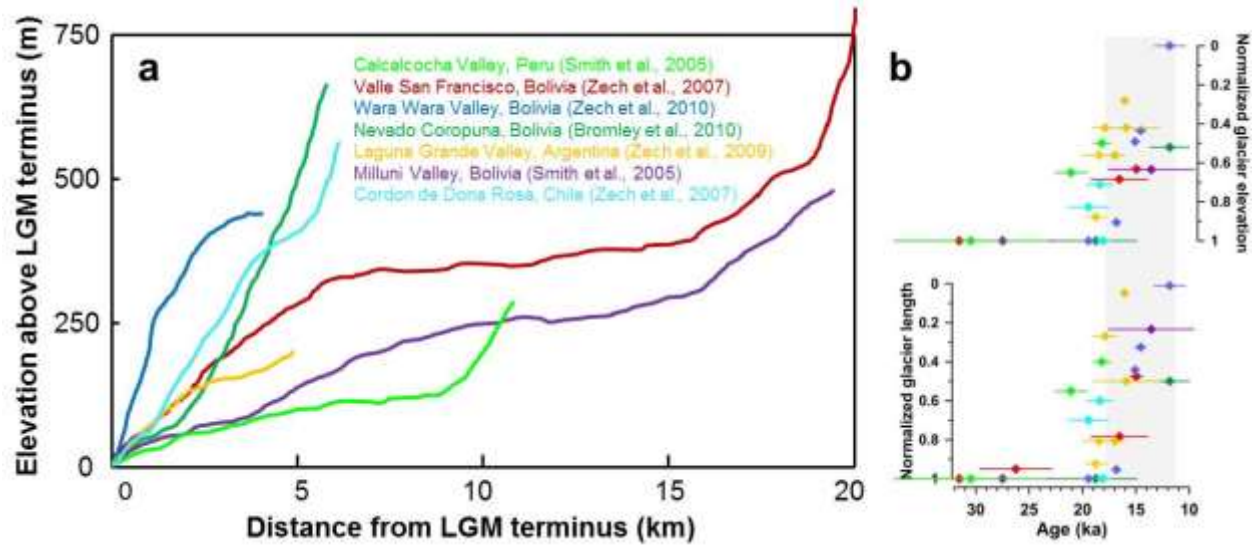
ICE simulation



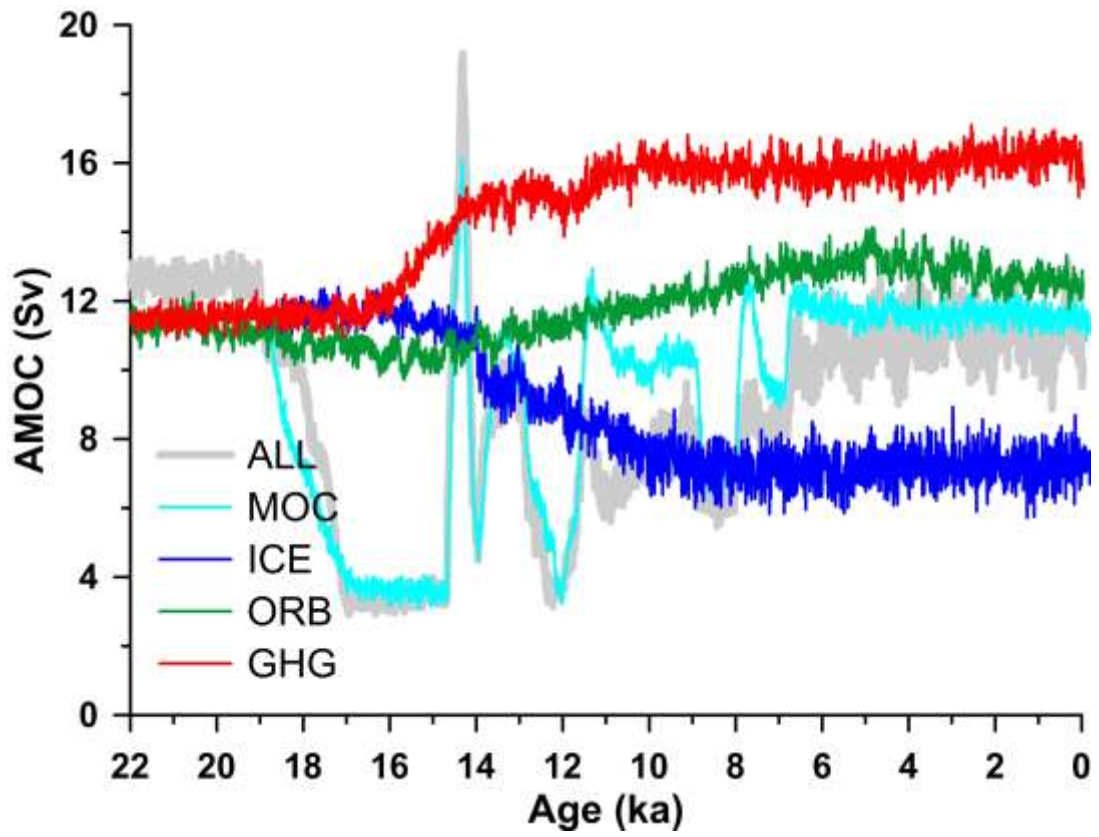
MOC simulation



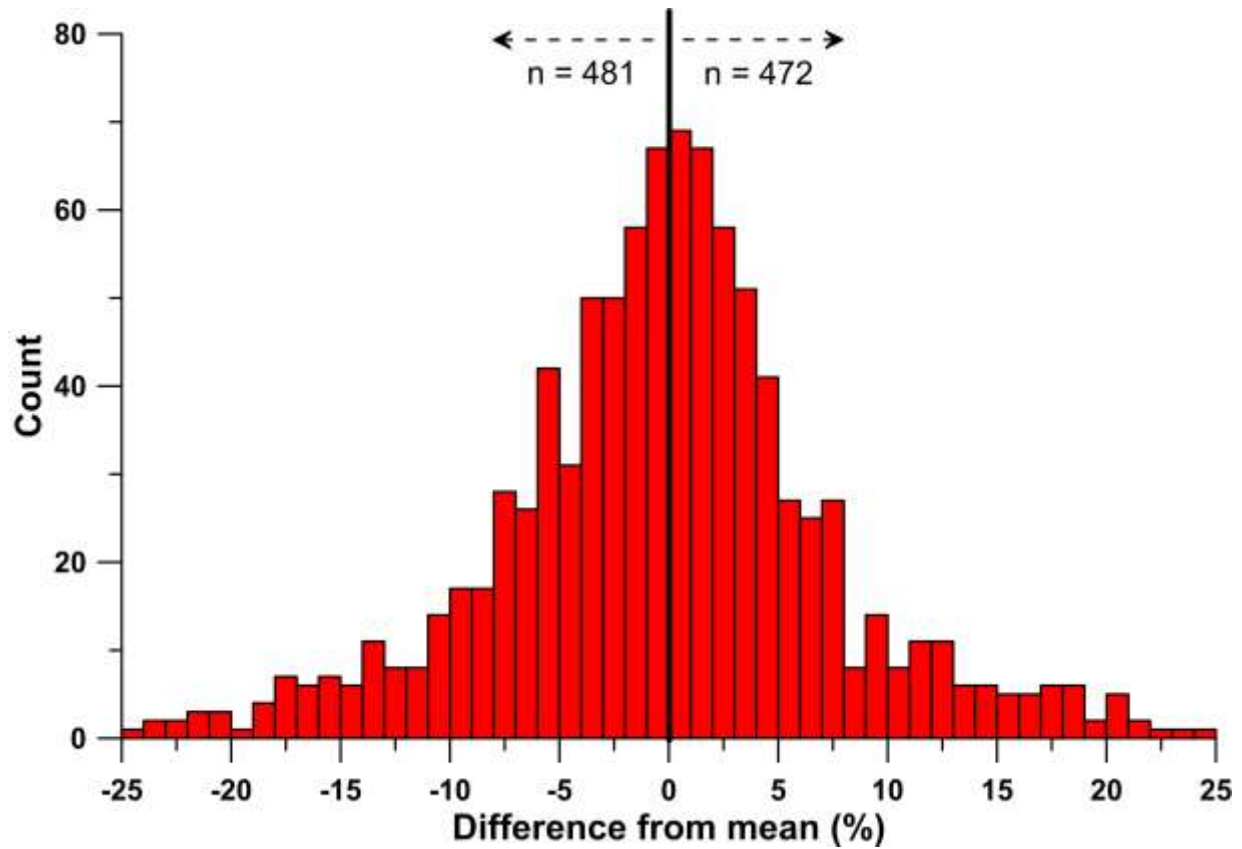
Supplementary Figure 4. Each figure above shows mean annual surface temperature at the Last Glacial Maximum (LGM; 21 ka for simulation ALL, GHG, and ORB, and 19 ka for simulation ICE and MOC) in the upper left panel for the various transient simulations, followed by surface temperature anomalies from the LGM at 18, 15, 12, 9, and 6 ka. Light blue dots on each map represent the locations of the glacial moraines synthesized in this study. Units are °C.



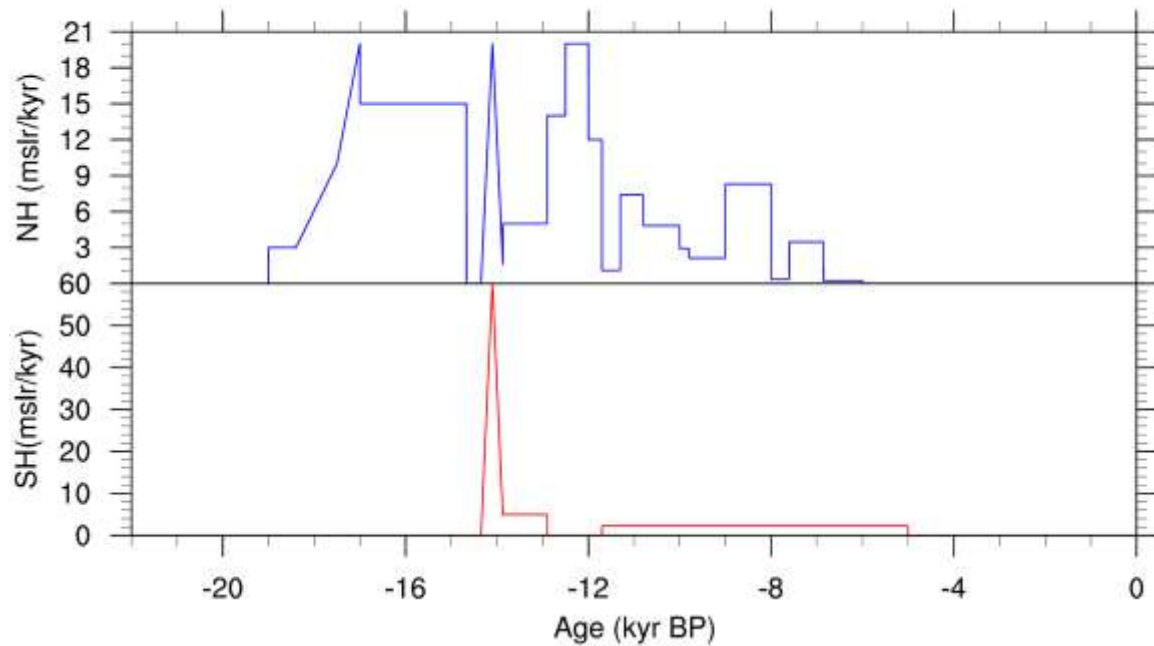
Supplementary Figure 5. (a) Valley profiles for several subtropical and equatorial Andes sites between the local LGM moraine and cirque headwall or modern glacier terminus. (b) The corresponding normalized glacier length and elevation histories for these sites. Gray bar highlights deglacial CO₂ and global temperature rise.



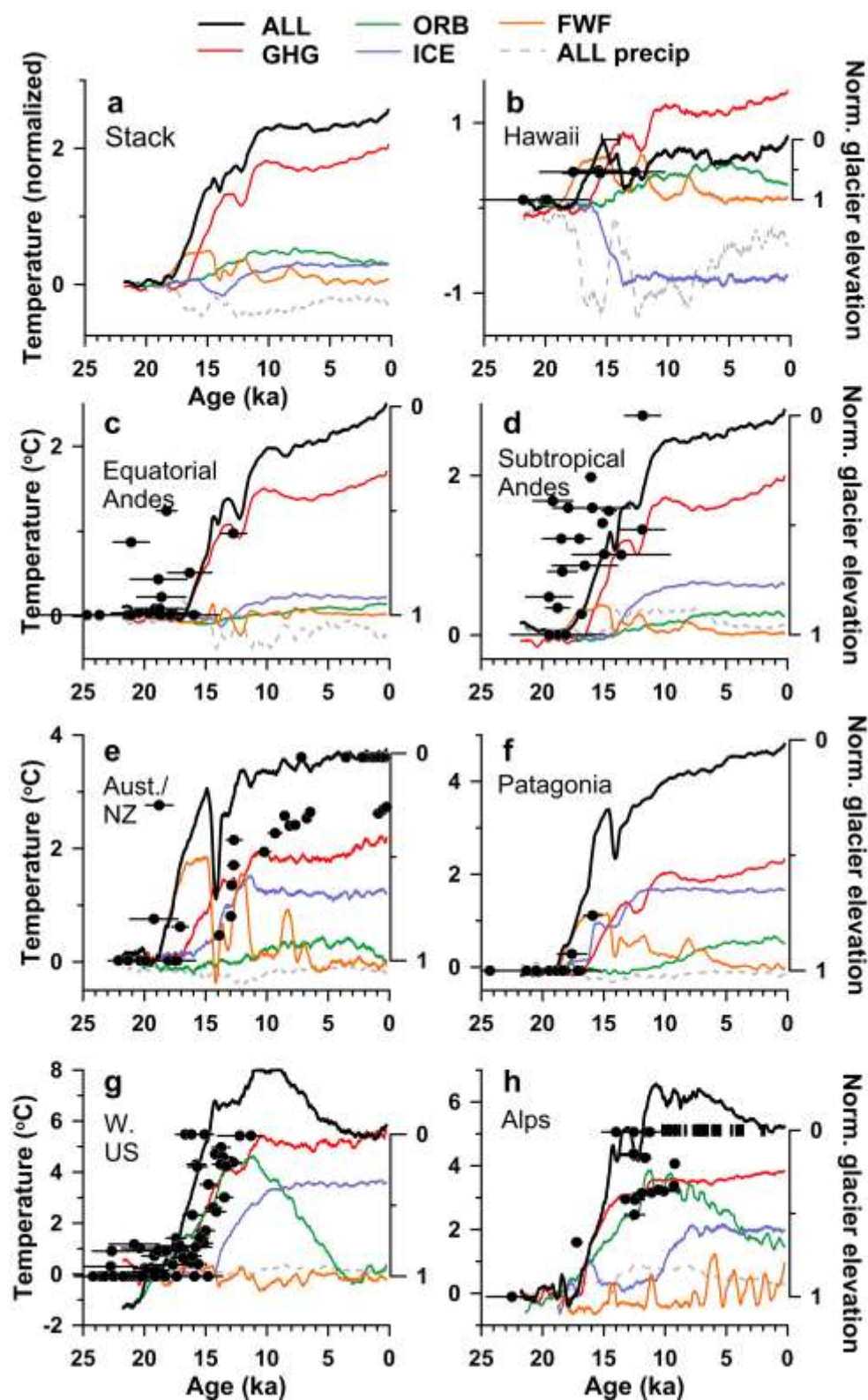
Supplementary Figure 6. Atlantic Meridional Overturning Circulation (AMOC) strength in the transient model simulations.



Supplementary Figure 7. Histogram showing the deviation of boulder ages from the mean age on each moraine. The number of samples on either side of the mean is indicated at the top of the figure. Only moraines with three or more boulder ages were included. Boulders considered outliers by the original authors were excluded.



Supplementary Figure 8. Northern and Southern Hemisphere freshwater forcing used in the MOC and ALL simulations.



Supplementary Figure 9. As in Figure 3 in the manuscript, but here showing normalized glacier elevation changes rather than length changes.

Supplementary Table 1

Sea-level, high-latitude (SLHL) production rates for ^{10}Be using nuclide-specific LSD scaling

Study	SLHL Production Rate (at $\text{g}^{-1} \text{y}^{-1}$)
Fenton et al. ⁴ – Norway	3.59 ± 0.18 (1.7)
Claude et al. ⁵ – Chironico landslide, Switzerland	3.94 ± 0.24 (2.1)
Ballantyne and Stone ⁶ – Scotland	4.01 ± 0.14 (0.2)
Kaplan et al. ⁷ – Puerto Bandera only, Patagonia	3.93 ± 0.10 (1.3)
Goehring et al. ⁸ – Norway	4.01 ± 0.12 (3.4)
Balco et al. ⁹ – NE North America	4.10 ± 0.20 (1.8)
Kelly et al. ¹⁰ – Peru *	4.13 ± 0.09 (2.5)
Putnam et al. ¹¹ – New Zealand	3.77 ± 0.08 (0.4)
Young et al. ¹² – Arctic	3.82 ± 0.13 (0.3)
Lifton et al. ¹³ – Lake Bonneville	4.20 ± 0.09 (1.1)
<i>Mean – All</i>	3.9 ± 0.2
<i>Mean – No Fenton⁺</i>	4.0 ± 0.1

Notes

Production rates only considered significant to one decimal place. 2nd decimal place only carried for rounding. Numbers in parentheses and italics following each site production rate are reduced χ^2 values for the fit at each site. Erosion rate and surface shielding corrections taken from original studies.

* Samples Q-47, Q-48 removed from dataset following Kelly et al.¹⁰. Samples Q-40a, Q-44a, and Q-83a also removed to prevent circularity since they were used to calculate surface erosion rates used for the other samples in the study. Including the latter 3 samples does not significantly change the site production rate (4.09 ± 0.09 ^{10}Be at $\text{g}^{-1} \text{y}^{-1}$). Assume erosion rates as stated in Kelly et al.¹⁰

+ Production rate from Fenton et al.⁴ (including all corrections from original study) fails Chauvenet's criterion¹⁴, and thus is considered an outlier for this dataset.

Supplementary Table 2

SLHL production rates for ^3He using nuclide-specific LSD scaling

Study	SLHL Production Rate (at $\text{g}^{-1} \text{y}^{-1}$)
Ackert et al. ¹⁵ – Patagonia*	127.6 ± 1.7 (19.3)
Amidon et al. ¹⁶ – Lake Bonneville (<i>pyroxene</i>)	125.1 ± 10.6 (0.2)
Blard et al. ¹⁷ – Mt. Etna*	131.4 ± 3.7 (15.7)
Blard et al. ¹⁷ – Hawaii*	149.9 ± 7.5 (1.6)
Blard et al. ¹⁸ – Altiplano	132.5 ± 4.1 (2.0)
Cerling and Craig ¹⁹ , Cerling ²⁰ – Western U.S. and France*	116.9 ± 3.5 (4.3)
Dunai and Wijbrans ²¹ – Canary Islands*	115.5 ± 13.1 (0.1)
Kurz et al. ²² – Hawaii*	134.6 ± 13.2 (7.4)
Goehring et al. ²³ – Lake Bonneville*	118.8 ± 2.1 (18.2)
Licciardi et al. ²⁴ – Western U.S.*	106.8 ± 4.3 (1.1)
Licciardi et al. ²⁵ – Iceland*	114.4 ± 5.3 (5.6)
Foeken et al. ²⁶ – Cape Verde Islands	94.7 ± 4.6 (2.5)
Mean	122 ± 14

Notes

Production rates only considered significant to whole numbers. First decimal place only carried for rounding. Numbers in parentheses and italics following each site production rate are reduced χ^2 values for the fit at each site. Erosion rate and surface shielding corrections taken from original studies

* Included in Goehring et al.²³

Supplementary Note 1

Excluding Asian data

Cosmogenic ages from Asia were excluded from our analysis for two reasons. First, moraine boulder ages from this region tend to exhibit relatively large scatter and there is thus greater uncertainty in the timing of glacier fluctuations there. For instance, Heyman⁴¹ synthesized 1855 ages on 113 moraines from the Tibetan Plateau and surrounding regions and found that more than half have uncertainties >15%, with many exceeding 30%. Second, most moraines dated in this region are from different valleys, and thus the retreat of individual glaciers from maximum to minimum extents cannot be scaled into the normalized units that we use to compare deglaciations across different regions and to potential forcings. For example, in Heyman's⁴¹ synthesis, only one valley has a 'Class A' moraine (those with uncertainties <15%) representing the local LGM (taken to be Marine Isotope Stage 2, or 30-18 ka) with an additional Class A deglacial-age moraine upvalley.

Supplementary Note 2

Greenhouse gas forcing

While we focus on the relationship between glacier retreat and CO₂ forcing in the main text, a more complete treatment of greenhouse-gas forcing would also include the effects from CH₄ and N₂O. Nonetheless, our approximation is justified since CO₂ contributed nearly 80% of the combined radiative forcing from these gases, and CO₂ forcing rose to within 10% of its preindustrial value by 11 ka.

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