

Supplementary Information for “The distribution of Earth’s wind-blown sand dunes”

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Supplementary Text 1

Many other secondary source maps were used to develop the map used in this study¹⁻²⁹. This text section provides further details on the manual mapping of dune presence.

Omissions

- The map is void where ERA5-Land reanalysis data is not provided, i.e., any 0.1°x0.1° grid cell where the area is majority water (except some lakes) and therefore omits many coastal dunes.
- Some dunes are too small (in height or planform) to observe at the mapping scale used.
- Some dunes are presently obscured by thick vegetation and cannot be discerned with the datasets used (e.g., COP30 is a Digital Surface Model, not a Digital Terrain Model). In some parts of Earth better datasets which would result in a more comprehensive map could be used (e.g., USGS 3DEP DTM), however, to ensure an unbiased map the same globally available datasets were used everywhere.
- Some aeolian dune fields can be identified but their surface expression is too modest, altered, or complex to identify a morphology, therefore they are omitted. This is particularly true for relict dunes. Some example regions are the Pampas, Kalahari, Low Qoz, Qoz Dango, Hausaland, Nara, East Caspian, Lapland, North Europe, North America, and Siberia.

New additions

This map includes dunes in the following regions which, to the best of the author’s knowledge, are not identified in other published maps including those used as secondary sources for this study:

- Nenetsia, Russia
- East of the Lut, Iran
- Afghanistan border with Tajikistan & Uzbekistan
- North-facing Himalaya, Tibet
- Sevrei, Mongolia
- Töv, Mongolia
- West Tabuk, Saudi Arabia
- Samara, Kazakhstan
- Jubaland coast, Somalia
- Senegal-Gambia border
- Between Bir Anzerane & Awsard, Western Sahara
- El Hank escarpment, Mauritania
- Downwind of Bel Bouna, Mauritania
- Djebel Ouarkiz, Algeria
- Between Taoundart & Amesmessa, Algeria
- Western Tazrouk Province, Algeria
- Western Tibetsi Mountains, Chad
- Around Sabkhat al Qunnayyin, Libya
- Between Ghurd as Sara & Al Jawf, Libya
- Eastern side of Jabal Nuqayy, Libya

Notes on definition of marginal dunes

Free elongating dunes are only classified if they have a sharp crest. Often there are mounds of sand in the lee of obstacles without a crest; we have not classified these as dunes. Similarly, we have not classified mounds of free sand without a clear wind-driven asymmetry in morphology as dunes. These, in some cases, may be dome dunes, however it is not possible to definitively distinguish them as aeolian.

Parabolic dunes have tremendous variety in morphology. This is partially because; they can form in unimodal winds with a variety of dispersions, have various stabilising forces (i.e., ice, vegetation, microbial mats, etc.), and, since they often are on the fringes of aeolian activity, can reactivate and degrade in unexpected ways on a variety of timescales from seasonal to Milankovitch. We have erred on a bullish identification of these dunes. We have, however, maintained a consistent approach to not classifying ‘blowouts’ without a clear downwind crest, i.e., an eroding sand patch surrounded by vegetation (or another stabilising force) containing no discernible bedforms, as dunes. These are prevalent in coastal settings and the Pampas.

Superimposed dunes with unique morphologies can prevail when either all morphologies are active, only some are active, or none. When none are active, and topography is degraded somewhat, these dunes can be difficult to identify. Therefore, there may be some locations where these are not included by mistake.

In some locations, particularly large ergs, there can be fields of consistently patterned dunes that, despite not displaying slipfaces, appear with context to be mature and free, akin to dome dunes but with high sand supply. We have classified these as dunes however they do not fit within a typical dune classification scheme. We have distinguished these from sand drifts which, despite having fine-scale micro-relief patch patterns, are not considered dunes, using historical Landsat imagery to assess their migration speed. A clear example of this distinction is seen between the Zaltan & Calanscio dune fields.

Sometimes there are clear aeolian patterns on surfaces (e.g., intermittently wet lake beds) that are not dunes. They can form for various reasons, such as through feedbacks between wind and vegetation or evaporation, but have not classified as dunes as they are not topographic forms.

There are certainly places which we have indicated dunes that would be up for expert debate, e.g., subtle relict nail dunes along the Warrego River, Australia, the crests of eroding hills south of the Nafusa Mountains, Libya, or slip faces developing upwind of the stoss of obstacles west of Mont Tousse, Chad. These surmount to a negligible fraction of the dataset ($<0.1\%$) so should not meaningfully impact the results.

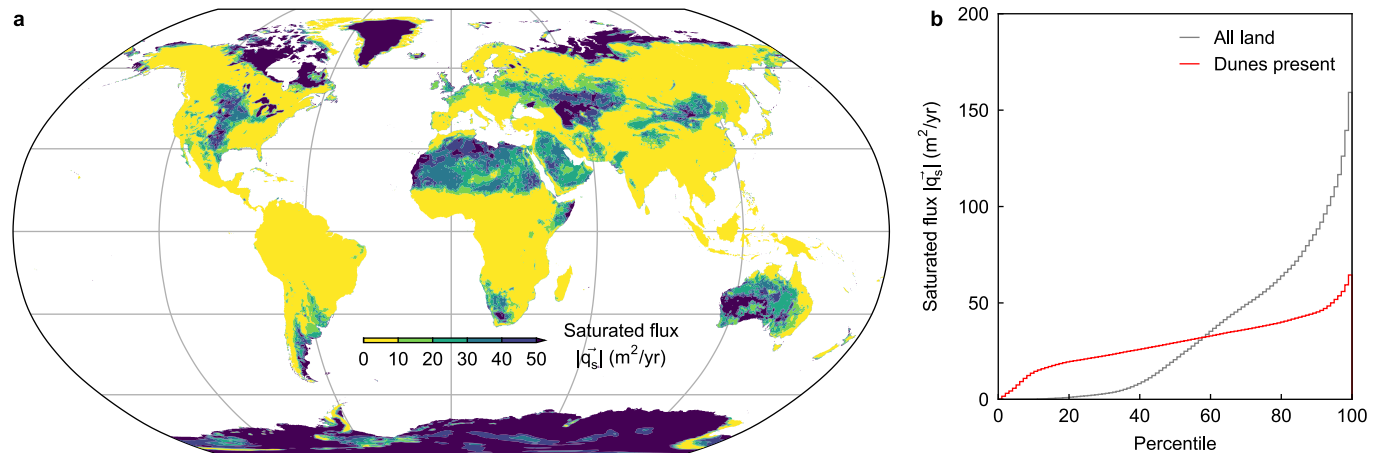
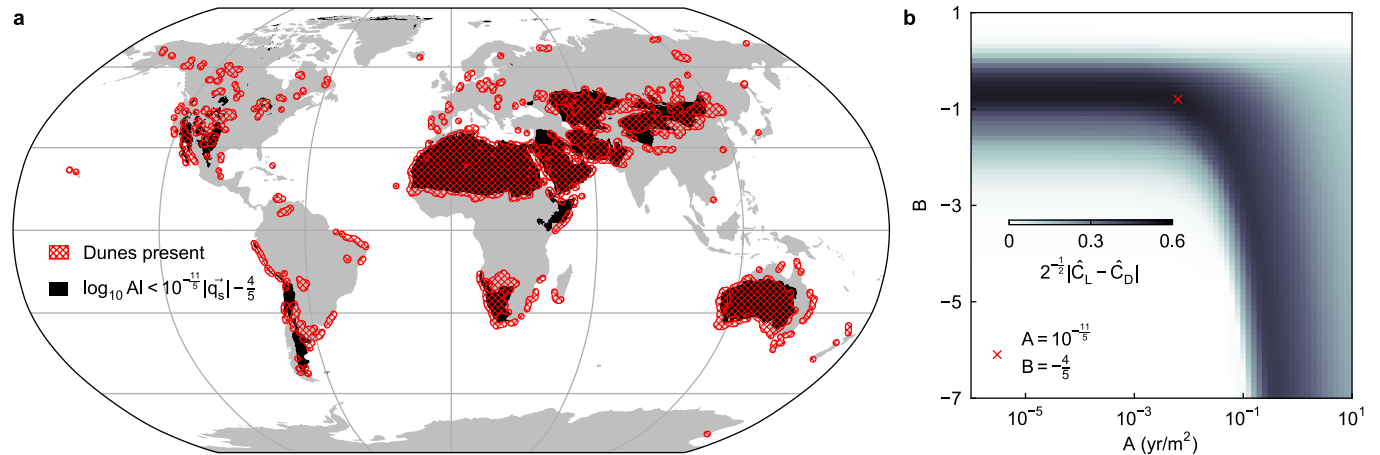
Supplementary Text 2

Past studies have introduced Mobility Indices for dune activity which have proved extremely useful for distilling the competing effects of wind and vegetation on dune activity³⁰⁻³⁷. These include myriad parameterisations for important hydroclimate and wind climate properties, such as the percentage of time wind exceed a threshold for sand transport^{30,32,34}, the mean wind speed raised to some power^{31,32}, the “RDP/DP”³⁵ (Methods M4), different flavours of the Aridity Index^{30,36}, and the ratio of actual to potential evapotranspiration³². Some also include reasonable memory effects which parameterise the influence of drought³² or existing non-equilibrium vegetation³⁷. Threshold values for these Mobility Indices are proposed which delineate when dunes become active or should be present³⁰⁻³⁴. Here we compare the precision and recall of the relation $\log_{10} AI < 10^{-11/5} |\bar{q}_s| - 4/5$ to commonly employed past criteria (Supplementary Table 1).

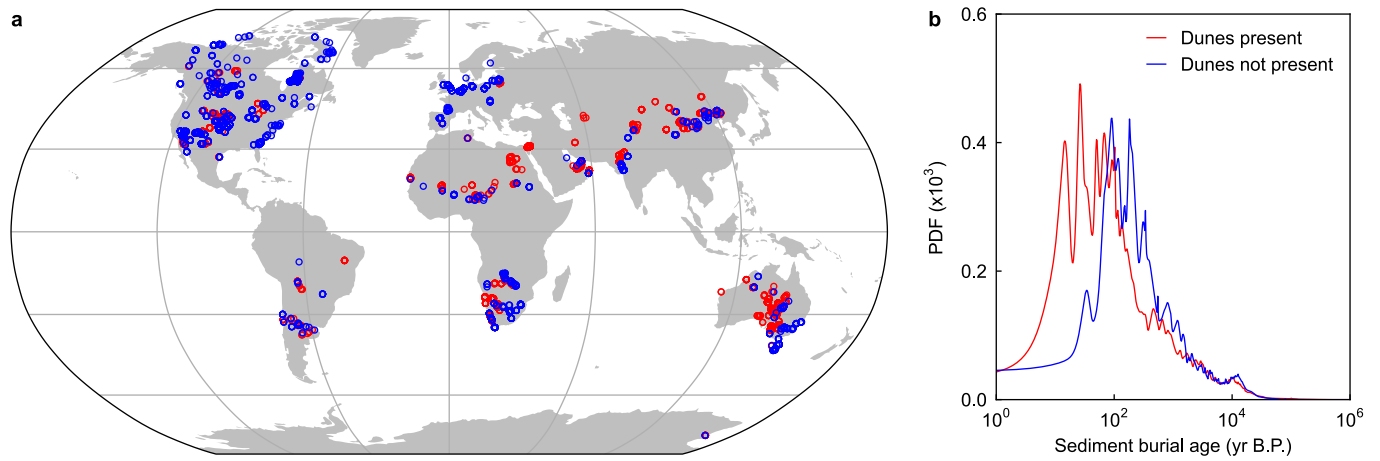
In order to make this comparison, we have made some minor assumptions. Lancaster³⁰ proposed that dune activity occurs when $W/AI > 50$, where W (%) is the percentage of time $|\bar{u}_{10}| > 4.5$ (m/s) and AI is the Aridity Index, where calculated in the Thornthwaite method—we assume the AI is consistent with the values used in this study. Ash & Wasson³² proposed that dune activity occurs when $|\bar{u}_{10}|^4/(AET/PET) > 10^4/3.8$ or $W^2/(AET/PET) > 10^4/5$, where AET is the actual evapotranspiration and PET is the potential evapotranspiration—we approximate AET/PET as AI . Thomas et al.³¹ proposed that dune activity occurs when $|\bar{u}_{10}|^3/AI > 70$, with a month-long memory effect due to drought—we neglect the memory effect.

Supplementary Table 1 shows that all other methods tested here have reduced precision relative to the relation in this study, apart from $|\bar{u}_{10}|^4/(AET/PET) > 10^4/3.8$ which has significantly worse recall. Note that the reduced recall relative to some other studies should be weighed against the increase in precision: since dune grid cells only account for 7.8% of the land surface, precision scores are far harder to increase relative to recall. Also note that, as discussed elsewhere in the study, all precisions are relatively low due to the role that other boundary conditions (i.e., flux divergence and sediment source location) play in the distribution of dunes, and recall is lower since relict dunes make up a small fraction of the dataset.

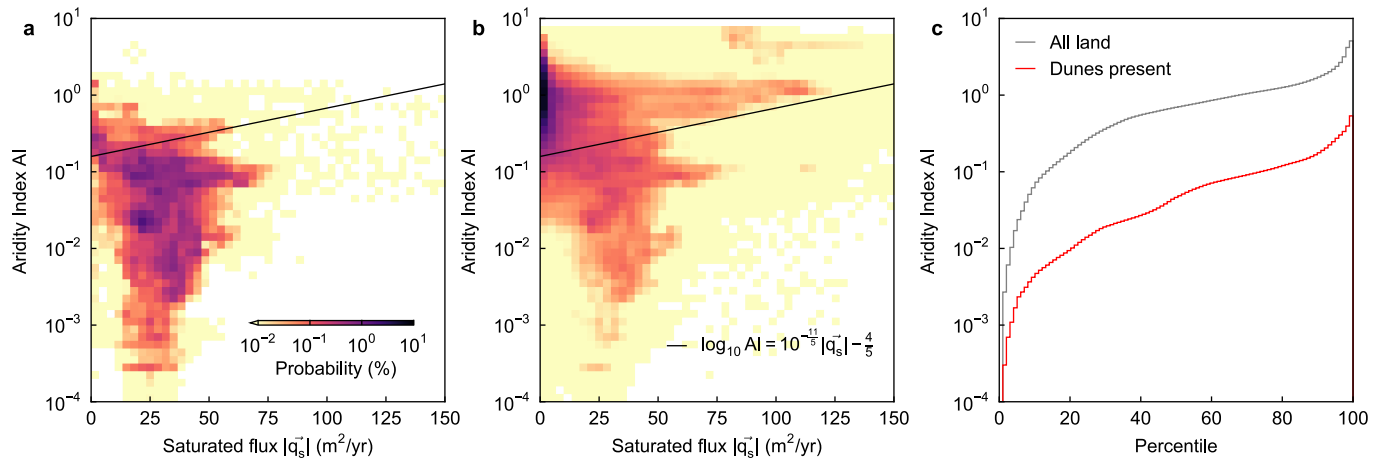
Supplementary Figures & Tables



Supplementary Fig. 2: Saturated sand flux magnitude. **a**, World map (ESRI:54030) of the average saturated sand flux magnitude $|\bar{q}_s|$ (m^2/yr). **b**, The area-weighted percentiles for land (grey) and regions where dunes are present (red).



Supplementary Fig. 3: Reported aeolian sediment burial ages. **a**, World map (ESRI:54030) of the locations of luminescence and radiocarbon dating burial ages in aeolian sediments from the INQUA Dune Atlas¹⁸ where errors are reported, where dunes with surface morphologies are present (red) or not (blue). **b**, The probability distribution functions, i.e., the simple-mean gaussian mixture, of all burial ages where dunes are present (red; N=2,295) or not (blue; N=2,097). Note that, in lieu of a detailed analysis of each specific age, this simplistic approach neglects contextual information about samples such as their depth or stratigraphic unit.



Supplementary Fig. 4: Likelihood of aridity and saturated flux. **a**, The joint area-weighted probability (log-scale) of Aridity Index AI and average saturated sand flux magnitude $|q_s^-|$ (m^2/yr) where dunes are present, and **b**, for all land, with the relation $\log_{10} AI = 10^{-11/5} |q_s^-| - 4/5$ (Methods M5) shown in black. **c**, The area-weighted percentiles for land (grey) and regions where dunes are present (red).

Study	Criterion	Precision	Recall
Lancaster (1988) ³⁰	$W/AI > 50$	26	96
Thomas et al. (2005) ³¹	$ u_{10}^- ^3/AI > 70$	23	98
Ash & Wasson (1983) ³²	$ u_{10}^- ^4/(AET/PET) > 10^4/3.8$	41	72
	$W^2/(AET/PET) > 10^4/5$	28	93
Muhs & Holliday (1995) ³³	$AI < 0.35$	25	98
Wasson (1984) ³⁴	$0.13W + \ln AI^{-1} > 1/0.21$	27	90
This study	$\log_{10} AI < 10^{-11/5} q_s^- - 4/5$	32	93

Supplementary Table 1: Precision and recall of mobility indices. Reported threshold criterion for dune mobility in other studies (Supplementary Text 2), and this study (Methods M5), and their area-weighted precision and recall scores (%).

References

- Zheng, Z., Zhang, X., Li, J., Ali, E., Yu, J., & Du, S. (2024). Global perspectives on sand dune patterns: Scale-adaptable classification using Landsat imagery and deep learning strategies. *ISPRS Journal of Photogrammetry and Remote Sensing*, 218, 781-801.
- Gunn, A., Casasanta, G., Di Liberto, L., Falcini, F., Lancaster, N., & Jerolmack, D. J. (2022). What sets aeolian dune height? *Nature Communications*, 13(1), 2401.
- Lu, Q., Gaur, M. K., & Squires, V. R. (Eds.). (2023). *Sand Dunes of the Northern Hemisphere: Distribution, Formation, Migration and Management, Volume 1*. CRC Press.
- Lancaster, N. (2013). *Geomorphology of desert dunes*. Routledge.
- Zhang, Z., Ma, P., La, Z., Zhang, Y., & Liang, A. (2023). Aeolian sediment provenance and transport in the upper and middle reaches of the Yarlung Zangbo River, Tibet Plateau. *Basin Research*, 35(2), 762-783.
- Liu, Y., Wang, Y. S., & Shen, T. (2019). Spatial distribution and formation mechanism of aeolian sand in the middle reaches of the Yarlung Zangbo River. *Journal of Mountain Science*, 16(9), 1987-2000.
- Dong, Z., Hu, G., Qian, G., Lu, J., Zhang, Z., Luo, W., & Lyu, P. (2017). High-altitude aeolian research on the Tibetan Plateau. *Reviews of Geophysics*, 55(4), 864-901.
- Doody, J. P. (Ed.). (1991). *Sand dune inventory of Europe*. JNCC/EUCC.
- Muhs, D. R. (2017). Evaluation of simple geochemical indicators of aeolian sand provenance: Late Quaternary dune fields of North America revisited. *Quaternary Science Reviews*, 171, 260-296.
- Lancaster, N., & Hesp, P. (Eds.). (2020). *Inland Dunes of North America*. Springer.
- Latrubesse, E. M., Stevaux, J. C., Cremon, E. H., May, J. H., Tatum, S. H., Hurtado, M. A., ... & Argollo, J. B. (2012). Late Quaternary megafans, fans and fluvio-aeolian interactions in the Bolivian Chaco, Tropical South America. *Palaeogeography, Palaeoclimatology, Palaeoecology*, 356, 75-88.
- Baas, A. C., & Delobel, L. A. (2022). Desert dunes transformed by end-of-century changes in wind climate. *Nature Climate Change*, 12(11), 999-1006.

13. Vimpere, L. (2024). Parabolic dune distribution, morphology and activity during the last 20 000 years: A global overview. *Earth Surface Processes and Landforms*, 49(1), 117-146.
14. Wolfe, S. (2025). Paleowind directions in northern North America from stabilized sand dunes [Dataset]. *Natural Resources Canada*.
15. Sayre, R. G., Comer, P., Hak, J., Josse, C., Bow, J., Warner, H., ... & Waruingi, L. (2013). A new map of standardized terrestrial ecosystems of Africa. *African Geographical Review*.
16. Wang, J., Wang, Y., Yan, C., & Qi, Y. (2013). 1: 100,000 desert (sand) distribution dataset in China. *Center NTPD, editor*.
17. Geoscience Australia (2006). GEODATA TOPO 250K Series 3 [Dataset]. *Geoscience Australia, Canberra*.
18. Lancaster, N., Wolfe, S., Thomas, D., Bristow, C., Bubenzer, O., Burrough, S., ... & Zárate, M. (2016). The INQUA dunes atlas chronologic database. *Quaternary International*, 410, 3-10.
19. Hesse, P. P. (2010). The Australian desert dunefields: formation and evolution in an old, flat, dry continent. *Geological Society, London, Special Publications*, 346(1), 141-164.
20. Miguel, L. L. A. J., & Castro, J. W. A. (2018). Aeolian dynamics of transgressive dunefields on the southern Mozambique coast, Africa. *Earth Surface Processes and Landforms*, 43(12), 2533-2546.
21. Warren, A. (1970). Dune trends and their implications in the central Sudan. *Z. Geomorphol. Suppl.*, 10, 154-180.
22. Crouvi, O., Schepanski, K., Amit, R., Gillespie, A. R., & Enzel, Y. (2012). Multiple dust sources in the Sahara Desert: The importance of sand dunes. *Geophysical Research Letters*, 39(13).
23. Hu, Z., Gao, X., Lei, J., & Zhou, N. (2021). Geomorphology of aeolian dunes in the western Sahara Desert. *Geomorphology*, 392, 107916.
24. Campos, J. C., & Brito, J. C. (2018). Mapping underrepresented land cover heterogeneity in arid regions: the Sahara-Sahel example. *ISPRS Journal of Photogrammetry and Remote Sensing*, 146, 211-220.
25. Maman, S., Porat, N., Meged, C., Blumberg, D. G., Tsoar, H., Mamedov, B., & Roskin, J. (2023). The Karakum and Kyzylkum Sand Seas. *Sand Dunes of the Northern Hemisphere: Distribution, Formation, Migration and Management: Volume 2: Characteristics, Dynamics and Provenance of Sand Dunes in the Northern Hemisphere*, 253.
26. Al-Dousari, A., Awadh, S., Bamousa, A., Al-Awadhi, J., Al-Dousari, N., Al-Rawi, M., ... & William, T. (2023). Barchan dunes properties in the Middle East. In *Sand Dunes of the Northern Hemisphere: Distribution, Formation, Migration and Management* (pp. 129-149). CRC Press.
27. Embabi, N. S. (2023). Barchan Dunes in Egypt Distribution and Formation. In *Sand Dunes of the Northern Hemisphere: Distribution, Formation, Migration and Management* (pp. 150-180). CRC Press.
28. Goudie, A. S., Goudie, A. M., & Viles, H. A. (2021). Dome dunes: Distribution and morphology. *Aeolian Research*, 51, 100713.
29. Abbasi, H. R., Opp, C., Groll, M., Rohipour, H., & Gohardoust, A. (2019). Assessment of the distribution and activity of dunes in Iran based on mobility indices and ground data. *Aeolian Research*, 41, 100539.
30. Lancaster, N. (1988). Development of linear dunes in the southwestern Kalahari, southern Africa. *Journal of Arid Environments*, 14(3), 233-244.
31. Thomas, D. S., Knight, M., & Wiggs, G. F. (2005). Remobilization of southern African desert dune systems by twenty-first century global warming. *Nature*, 435(7046), 1218-1221.
32. Ash, J. E., & Wasson, R. J. (1983). Vegetation and sand mobility in the Australian desert dunefield. *Zeitschrift fur Geomorphologie*, 45(Supp.), 7-25.
33. Muhs, D. R., & Holliday, V. T. (1995). Evidence of active dune sand on the Great Plains in the 19th century from accounts of early explorers. *Quaternary Research*, 43(2), 198-208.
34. Wasson, R. J. (1984). Late Quaternary palaeoenvironments in the desert dunefields of Australia. In *Late Cainozoic palaeoclimates of the Southern Hemisphere. International symposium held by the South African Society for Quaternary Research; Swaziland* (pp. 419-432).
35. Tsoar, H. (2005). Sand dunes mobility and stability in relation to climate. *Physica A: Statistical Mechanics and its Applications*, 357(1), 50-56.
36. Talbot, M. (1984). Late Pleistocene rainfall and dune building in the Sahel. In *Palaeoecology of Africa and the surrounding Islands*, volume 16 (pp. 203-214).
37. Yizhaq, H., Ashkenazy, Y., & Tsoar, H. (2007). Why do active and stabilized dunes coexist under the same climatic conditions? *Physical Review Letters*, 98(18), 188001.